



# New insights into the reproductive cycle of two Great Scallop populations in Brittany (France) using a DEB modelling approach

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## ABSTRACT

The present study aimed to improve understanding of the environmental conditions influencing the reproductive cycle of the great scallop *Pecten maximus* in two locations in Brittany (France). We also evaluated potential consequences of future climate change for reproductive success in each site.

We simulated reproductive traits (spawning occurrences and synchronicity among individuals) of *P. maximus*, using an existing Dynamic Energy Budget (DEB) model. To validate and test the model, we used biological and environmental datasets available for the Bay of Brest (West Brittany, France) between 1998 and 2003. We also applied the scallop DEB model in the Bay of Saint-Brieuc (North Brittany, France) for the same period (1998–2003) to compare the reproductive cycle in different environmental conditions. In order to accurately model the *P. maximus* reproductive cycle we improved the scallop DEB model in two ways: through (1) energy acquisition, by incorporating microphytobenthos as a new food source; and (2) the reproductive process, by adding a new state variable dedicated to the gamete production. Finally, we explored the effects of two contrasting IPCC climate scenarios (RCP2.6 and RCP8.5) on the reproductive cycle of *P. maximus* in these two areas at the 2100 horizon.

In the Bay of Brest, the simulated reproductive cycle was in agreement with field observations. The model reproduced three main spawning events every year (between May and September) and asynchronicity in the timing of spawning between individuals. In the Bay of Saint-Brieuc, only two summer spawning events (in July and August) were simulated, with a higher synchronicity between individuals. Environmental conditions (temperature and food sources) were sufficient to explain this well-known geographic difference in the reproductive strategy of *P. maximus*. Regarding the forecasting approach, the model showed that, under a warm scenario (RCP8.5), autumnal spawning would be enhanced at the 2100 horizon with an increase of seawater temperature in the Bay of Brest, whatever the food source conditions. In the Bay of Saint-Brieuc, warmer temperatures may impact reproductive phenology through an earlier onset of spawning by 20 to 44 days depending on the year.

## 1. Introduction

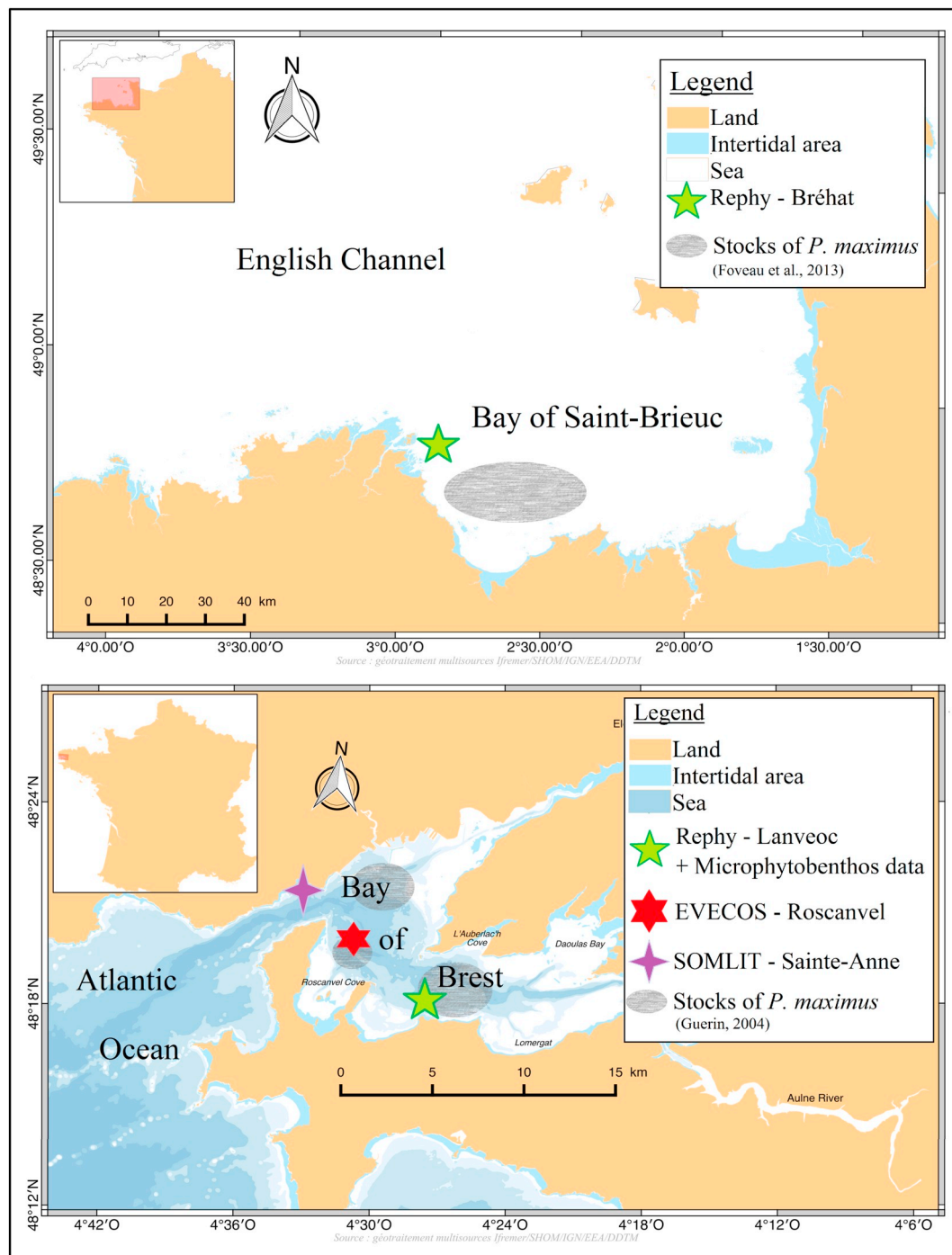
The great scallop, *Pecten maximus* (Linnaeus, 1758) inhabits many sublittoral environments along Northeast Atlantic coasts from northern Norway to the Iberian Peninsula (Ansell, 1991). In France, the species is particularly abundant along the coast of northern Brittany, where it sustains one of the most important French commercial fisheries both in terms of landings and of socio-economic value (> 300 fishing boats; ICES, 2015). The main fishing areas are located in the Bay of Brest, connected to the Iroise Sea, and in the

Bay of Saint-Brieuc, open to the English Channel (Fig. 1). While some of the highest scallop densities are found in the Bay of Saint-Brieuc, in part due to sustainable exploitation measures, the scallop stock in the Bay of Brest is lower and highly dependent on hatchery produced spat since 1983.

From a biological point of view, scallops, like most other bivalves, are filter feeders and consume phytoplankton. However, since they live settled into the surface layer of the bottom, they are also thought to use the epibenthic layer as an important food source (see review in Shumway, 1990). Concerning the reproductive cycle, *P. maximus* is a

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**Fig. 1.** Maps of the two study sites, the Bay of Saint-Brieuc and the Bay of Brest, showing the position of the bimonthly great scallop monitoring area of (EVECOS) and the three environmental monitoring sites: the REPHY stations at Lanvéoc and Bréhat and the SOMLIT station at Sainte-Anne.

functional hermaphrodite species, it has a pelagic larval stage during approximately 1 month after fertilization, switching to a benthic life after metamorphosis. Its reproductive strategy is more surprising as its spawning period varies according to the geographical location of the population (see review by Gosling (2004)). There can be between one major summer spawning and more than three spawnings in the period from spring to early autumn. At a small regional scale, geographical differences can be very marked: scallops from the Bay of Brest show low inter-individual synchronism, with multiple partial spawnings from early spring to autumn and almost no resting stage, whereas the population from the Bay of St-Brieuc is almost synchronous, with one or

two major spawnings over a short period (July–August), with a sexual rest stage then observed in autumn and winter (e.g. Devauchelle and Mingant, 1991; Paulet et al., 1997).

A major part of this phenotypic variability has been attributed to differences in environmental conditions such as food sources, temperature and photoperiod, which are known to influence gametogenesis and fecundity in marine invertebrates. For example, Clereboudt and Himmelman (1996) showed that an increase in temperature and food availability increased reproductive effort in *Placopecten magellanicus*. In *P. maximus*, quantity and quality of food sources also have an impact on hatching rate (Soudant et al., 1996), and laboratory experiments

showed that spring conditions (regular increase of temperature and photoperiod) favoured gonad growth, whereas winter conditions (regular decrease of temperature and daylight duration) were associated with somatic growth of the adductor muscle and digestive gland (Saout et al., 1999; Lorrain et al., 2002). More recently, Chauvaud et al. (2012) and Lavaud et al. (2014) have proposed complementary approaches to quantitatively evaluate effects on environmental factors on growth and reproduction of scallops. However, the relative importance of these variables remains difficult to disentangle, especially under natural conditions.

Climate models and observations to date indicate that the Earth will warm between two (IPCC scenario RCP2.6) and six degrees Celsius (IPCC scenario RCP8.5) over the next century, depending on how fast carbon dioxide emissions increase. The ocean absorbs most of this excess heat, leading to rising seawater temperatures (e.g., IPCC, 2014; Appendix A). Increasing ocean temperatures will deeply affect marine species and ecosystems. Understanding the potential effects of climate change on the timing of life-history events such as the onset of gametogenesis, spawning, hatching and larval metamorphosis is important for benthic ecology but also for aquaculture and fisheries production. The phenology of these key life-history events has been investigated in several ecosystems and in many species (e.g., Beukema et al., 2009; Menge et al., 2009; Shephard et al., 2010; Valdizan et al., 2011; Morgan et al., 2013), although these studies often had limited spatial and/or temporal resolution. Mechanistic modelling provides a complementary tool to analyse climate effects on life-history traits, identify interactions between multiple stressors, and make predictions about future condition scenarios at a larger spatiotemporal scale. In recent decades, bivalve growth and reproduction have been successfully modelled (e.g. Bernard et al., 2011; Thomas et al., 2016; Montalto et al., 2017; Gourault et al., 2018, this issue) using mechanistic models based on Dynamic Energy Budget theory (DEB; Kooijman, 2010). This theory makes it possible to quantify the energy flows within an individual from ingestion to maintenance, growth, development, and reproduction in relation to environmental conditions.

In this context, the present study aims to improve understanding of the environmental factors influencing the reproductive cycle of *P. maximus* using a DEB model and the potential effects of climate change on the reproductive activity of this species. Our work is based on an existing DEB model developed for the great scallop in the Bay of Brest (Lavaud et al., 2014) that we then improved by adding detail on the reproductive processes. To evaluate the ability of the model to simulate reproductive processes under various conditions, we tested it over six years in the Bay of Brest (1998–2003) and in the two main locations hosting scallop populations in Brittany: the Bay of Brest and the Bay of Saint-Brieuc. In a second step, using two IPCC climate scenarios at the 2100 horizon, we examined the potential consequences of future climate change on the reproductive activity of this emblematic species in each of the two sites.

## 2. Material and methods

### 2.1. Study sites

The Bay of Brest is a semi-enclosed coastal ecosystem located in western Brittany, France, connected to the Atlantic Ocean by a deep narrow strait. The bay itself covers an area of nearly 180 km<sup>2</sup>, with an average depth of 8 m. Two rivers flow into the bay: the Elorn (watershed of 402 km<sup>2</sup>) and the Aulne (watershed of 1842 km<sup>2</sup>) (Fig. 1). Temperature and phytoplankton concentration are continuously monitored at two locations in the Bay: Lanvéoc station in the southern part of the Bay (data provided by the REPHY network – Phytoplankton and Phycotoxin monitoring NETwork, Ifremer, e.g. Belin et al., 2017) and Sainte-Anne station in the north-western part

(data provided by the SOMLIT – “Service d’Observation en Milieu Littoral”, INSU-CNRS, Brest). Lanvéoc station (48°29′ N, 04°46′ W; Fig. 1) has a depth range of 6–9 m at lowest spring tides and a bottom composed of sandy mud, with broken shells and pebbles. Sainte-Anne station is located at the entrance to the Bay of Brest (48°21′ N, 04°33′ W; Fig. 1).

The Bay of Saint-Brieuc is located in northern Brittany (France), 150 km from the Bay of Brest (48°32′ N, 02°40′ W; Fig. 1), in the western part of the English Channel. This bay of 800 km<sup>2</sup> harbours a large wild scallop population in its inshore shallow waters (≤ 30 m). It is subject to an extreme tidal regime with a tidal range between 4 m at neap tides and nearly 13 m during spring tides. Seawater temperature and phytoplankton concentration are monitored at the Bréhat station located in the western part of the bay (Fig. 1).

### 2.2. Scallop biological data

Scallop growth and reproduction were monitored from 1998 to 2003 at Roscanvel in the west of the Bay of Brest (48°20′ N, 04°30′ W; Fig. 1). This site, known to host the highest density of *P. maximus* in the bay, is characterized by mixed sandy and silty sediments. It was integrated into a multi-annual monitoring network (EVECOS database provided by the “Observatoire Marin de l’IUEM, Brest, France”). A sample of 20 adult scallops (3 years old) was collected by dredging on a biweekly to monthly basis in 30-m deep waters. The scallops were brought back to the laboratory where the muscle, gonads and digestive gland were dissected out. Total wet flesh mass and total dry flesh mass (DFM) of each organ were measured for each individual. In order to compare masses obtained for different sized scallops, dry mass was standardized following the formula of Bayne et al. (1987):

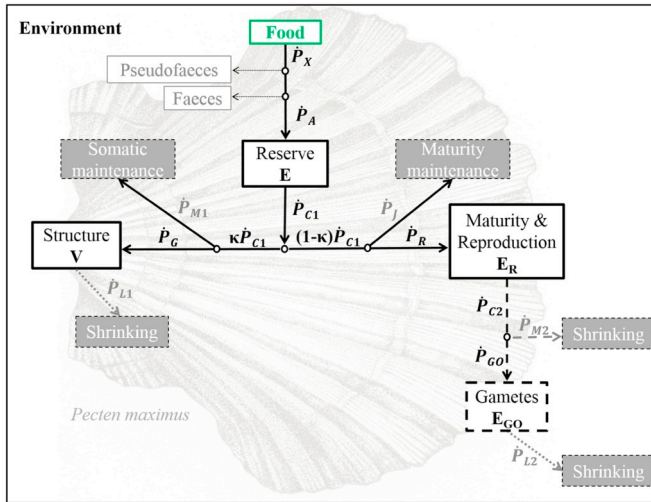
$$W_r = \left( \frac{L_r}{L_m} \right)^3 W_m$$

where  $W_r$  is the standardized mass of an individual of standard shell height  $L_r$  and  $W_m$  is the measured mass of an individual of measured shell height  $L_m$ . Length and mean daily shell growth rate (DSGR) were measured according to the method proposed by Chauvaud et al. (2012) (see Lavaud et al., 2014 for more detailed information on these data).

Additionally, four additional *P. maximus* reproductive cycle traits were recorded through EVECOS monitoring: the onset of gametogenesis, the number and timing of each main spawning within the reproductive season and the reproductive investment (DFM difference before and after spawning).

### 2.3. The scallop DEB model

The scallop DEB model was derived from the standard DEB model described by Kooijman (2010) and first applied to *P. maximus* by Lavaud et al. (2014). The DEB model is a mechanistic model that describes the dynamics of three state variables: E, the energy in reserve, V, the volume of structure, and  $E_R$ , the energy allocated to development and reproduction. To improve the accuracy with which DEB models model reproductive activity, Bernard et al. (2011) refined the processes of energy allocation to gametogenesis and resorption in the model, such that a fourth state variable,  $E_{GO}$ , the energy in gametes, was added to the existing scallop DEB model (Fig. 2). Briefly, the model can be explained as follows: the reserve mobilization rate,  $\dot{p}_{C1}$ , is divided into two parts. A first constant fraction,  $\alpha$ , is allocated to structural growth and maintenance and the remainder,  $1-\alpha$ , is allocated to development (in juveniles), reproduction (in adults) and maturity maintenance. Energy allocation to gonad construction is modelled through the gamete mobilization rate,  $\dot{p}_{C2}$ . Priority in energy allocation is always given to



**Fig. 2.** Schema of the *P. maximus* DEB model with four state variables adapted from Bernard et al., 2011 (Mod-2). Sources of energy to pay for somatic maintenance during prolonged starvation are indicated by grey dotted arrows. Modifications of the standard DEB model (3 state variables; Mod-1) are represented by dashed arrows for  $\dot{P}_{C2}$ ,  $\dot{P}_{GO}$ ,  $\dot{P}_{M2}$  and  $\dot{P}_{L2}$ .

maintenance costs:  $\dot{p}_{M1}$  for maturity maintenance and  $\dot{p}_j$  for somatic maintenance. During starvation periods, the gametogenesis flux is reallocated to somatic and maturity maintenance through secondary maintenance,  $\dot{p}_{M2}$ . If  $\dot{p}_{M2}$  does not provide enough energy to cover all maintenance costs, the gamete resorption rate,  $\dot{p}_{L2}$ , is activated. In case of extreme starvation, structure can be broken down at the rate  $\dot{p}_{L1}$ . The corresponding set of equations can be found in Gourault et al. (2018, this issue).

Regarding food assimilation, a classical scaled functional response (Holling type II) is generally calculated in the model (Kooijman, 2010), using one food source (for bivalves, this essentially consists of phytoplankton cells). However, many studies focusing on modelling the energy dynamics of filter feeders have shown the need and benefit of adding a second food source to improve the food proxy (Alunno-Bruscia et al., 2011; Bernard et al., 2011; Saraiva et al., 2011). Lavaud et al. (2014) included the Synthesizing Units (SUs) concept (Kooijman, 2010; Saraiva et al., 2011) into the scallop DEB model to consider selection of particles based on their size and/or quality. The equations for the SU concept can be found in Lavaud et al. (2014).

In this study we compared the previous model of Lavaud et al. (2014), hereafter referred to as “Mod-1”, with our DEB model (with the extra state variable  $E_{GO}$ ), hereafter referred to as “Mod-2” (Table 1). Two versions of the Mod-2 model were used in order to test different food sources in the model: (1) phytoplankton as a first food proxy and particulate organic matter (POM) as a second food proxy (Mod-2A) and (2) a mix of microphytobenthos and phytoplankton as a first food proxy and POM as a second food proxy (Mod-2B). All the model parameters are given in Table 2. Simulations were performed using R software (3.3.3 version).

**Table 1**  
The three scallop DEB models tested in this study.

	Mod-1 (Lavaud et al., 2014)	Mod-2A	Mod-2B
State variables	V, E, $E_R$	V, E, $E_R$ , $E_{GO}$	V, E, $E_R$ , $E_{GO}$
X-type food	Phytoplankton <sup>b</sup>	Phytoplankton <sup>b</sup>	Phytoplankton <sup>b</sup> + Microphytobenthos <sup>a</sup>
Y-type food	POM <sup>a</sup>	POM <sup>a</sup>	POM <sup>a</sup>

<sup>a</sup> data only available for the Bay of Brest.

<sup>b</sup> data available for the Bay of Brest and Bay of Saint-Brieuc

## 2.4. Model calibration

The model was calibrated with field data observed over the 1998–2003 period in the Bay of Brest. Model parameters (Table 2) were mostly taken from Lavaud et al. (2014), but some parameters were recalibrated for this study. First, we set a new value for the ultimate shell length  $L_{w\infty}$  (i.e., the maximum observed length reached in optimal condition i.e.  $f = 1$ ) at 20 cm instead of 12 cm. Some field studies have shown that adult scallops can reach 16 cm in the most favourable conditions (Mason, 1957; Chauvaud et al., 2012), so we set  $L_{\infty}$  above this value. According to DEB theory,  $L_{\infty}$  is calculated through the following equation:

$$L_{w\infty} = f \frac{L_m}{\delta_M} = f \frac{\left( \frac{\kappa [\dot{p}_{Am}]}{[\dot{p}_M]} \right)}{\delta_M}$$

where  $\{\dot{p}_{Am}\}$  is the maximum surface specific assimilation,  $[\dot{p}_M]$  is the volume-specific maintenance costs,  $\kappa$  is the allocation fraction to growth and maintenance and  $\delta_M$  is the shape coefficient. We modified the values of  $\kappa$ ,  $\{\dot{p}_{Am}\}$  and  $[\dot{p}_M]$ , while keeping  $\delta_M = 0.36$ . We estimated the values of  $\{\dot{p}_{Am}\}$  from Strohmeier et al. (2009) and a known value of  $[\dot{p}_M]$  at the same reference temperature (Emmery, 2008). Therefore, we were able to recalculate  $\kappa = 0.38$ .

To account for variability in the initial conditions between individuals, we simulated 20 individuals in each scenario (i.e., 20 different individual growth trajectories) by setting 20 different initial conditions of size and weight (i.e., first sampling of the year from EVECOS monitoring). Initial values for the four state variables ( $E$ ,  $V$ ,  $E_R$  and  $E_{GO}$ ) were calculated using the equations given in Table 3 from the measurements obtained in the first sampling of the year. Individual growth simulations were then pooled together to compute average growth patterns and standard deviation.

Three parameters control spawning in our model: the gonado-somatic ratio GSI, photoperiod and phytoplankton concentration. Threshold values for each of these three parameters were set as follows: GSI = 15% (estimated according to biological data from EVECOS monitoring), photoperiod (Photo) = 14 h (spawning is possible only if the daylength is above 14 h; Saout et al., 1999) and a phytoplankton concentration threshold (Phyto) =  $2.50 \cdot 10^5 \text{ cell L}^{-1}$  (average value corresponding to the beginning of a spring bloom; Paulet et al., 1997). In contrast to Lavaud et al. (2014), we calculated the GSI as the ratio between dry gonad mass and DFM, rather than as the ratio between wet gonad weight and cubic length. To assess the reproductive effort, individual DFM loss was estimated as the difference between individual DFM before and after spawning. Because spawning is mostly partial in *P. maximus*, 85% of the energy stored in  $E_{GO}$  was released as gametes at spawning and the remaining 15% was kept in the buffer for a potential subsequent spawning if environmental conditions remained optimal until winter. If conditions deteriorated, energy stored in the reproduction buffer was then used for the maintenance.

Field studies conducted in the Bay of Saint-Brieuc in the 1980s (e.g. Paulet et al., 1988) showed that phytoplankton blooms were much lower in this bay compared with the Bay of Brest. Over the 1998–2003 period, maximum phytoplankton concentrations in the Bay of Saint-Brieuc were always below the phytoplankton



**Table 2**List of the parameters implemented in the scallop DEB model. All rate parameters are given at  $T_1 = 15^\circ\text{C}$  ( $= 288.15\text{ K}$ ).

Description	Symbol	Value	Units	Reference
Biological parameters				
Shape coefficient	$\delta_M$	0.36	–	Lavaud et al. (2014)
Length at puberty (reproductive maturity)	$L_p$	4	cm	Lavaud et al. (2014)
Food assimilation				
Radius of X-type particle	$r_X$	15	$\mu\text{m}$	Lavaud et al. (2014)
Radius of Y-type particle	$r_Y$	15	$\mu\text{m}$	Lavaud et al. (2014)
Yield of reserve on X-type particle	$Y_{EX}$	0.7	–	Lavaud et al. (2014)
Yield of reserve on Y-type particle	$Y_{EY}$	0.4	–	Lavaud et al. (2014)
Max. specific filtration rate of X-type particle	$\dot{F}_{Xm}$	6	$\text{J d}^{-1} \text{cm}^2$	this study
Max. specific filtration rate of Y-type particle	$\dot{F}_{Ym}$	2	$\text{J d}^{-1} \text{cm}^2$	Lavaud et al. (2014)
Mol.weight of X-type particle	$w_X$	26.95	$\text{g mol}^{-1}$	Lavaud et al. (2014)
Mol.weight of Y-type particle	$w_Y$	25.4	$\text{g mol}^{-1}$	Lavaud et al. (2014)
Reserve parameters				
Volume-specific maintenance costs	$[\dot{p}_M]$	24	$\text{J cm}^{-3} \text{d}^{-1}$	Emmery (2008)
Energy conductance	$\dot{v}$	0.183	$\text{cm d}^{-1}$	van der Meer (2006)
Energy content of 1 g (dry weight) of reserve	$\rho_E$	19,849	$\text{J g}^{-1}$	Lavaud et al. (2014)
Molecular weight of reserve	$w_E$	23.9	$\text{g mol}^{-1}$	Lavaud et al. (2014)
Structure parameters				
Volume specific cost of growth	$[E_G]$	2959	$\text{J cm}^{-3}$	Lavaud et al. (2014)
Allocation fraction to growth and maintenance	$\kappa$	0.38	–	this study
Density of structure	$d_V$	0.12	$\text{g cm}^{-3}$	Lavaud et al. (2014)
Energy content of 1 g (dry weight) of structure	$\rho_V$	19,849	$\text{J g}^{-1}$	Lavaud et al. (2014)
Yield of structure tissue used for maintenance	$Y_{L1}$	1	–	Bernard et al. (2011)
Reproduction parameters				
Reproduction efficiency	$\kappa_{Go}$	0.70	–	this study
Density of gonad	$d_{Go}$	0.276	$\text{g cm}^{-3}$	this study
Yield of gonad tissue used for maintenance	$Y_{L2}$	1	–	Bernard et al. (2011)
Energy content of 1 g of gonad	$\rho_{Go}$	21,630	$\text{J g}^{-1}$	Bernard et al. (2011)
Temperature threshold for spawning	$T_S$	16	$^\circ\text{C}$	this study
Gonado-somatic index threshold for spawning	$GSI$	0.15	–	this study
Temperature effect				
Arrhenius temperature	$T_A T_A$	8990	K	Lavaud et al. (2014)
Lower boundary tolerance range	$T_L$	273.15	K	Lavaud et al. (2014)
Arrhenius temperature for lower boundary	$T_{AL}$	50,000	K	Lavaud et al. (2014)

**Table 3**

Initial state variables of the scallop DEB model. Values vary for each of the 20 individuals simulated, according to their initial length and initial dry flesh mass.

Initial conditions	Symbol	Equation	Units
Initial length	$L_0$	Observed measurements in the first sampling of the year	cm
Initial dry flesh mass	$W_{d0}$		g
Initial structure	$V_0$	$= (\delta_M L_0)^3$	$\text{cm}^3$
Initial reserve	$E_0$	$= f[E_m] V_0$	J
Initial gametes	$E_{GO0}$	$= \frac{(W_{d0} \times \rho_{GO})}{\kappa_{GO}}$	J
Initial reproduction	$E_{R0}$	$= \rho_E(W_{d0} - V_0 d_V) - E_0 - E_{GO0}$	J

concentration threshold set for Bay of Brest. Therefore, we hypothesised that phytoplankton concentration might not be relevant for triggering spawning in this more oligotrophic bay. Consequently, we added a temperature criterion based on the findings of Fifi (2004), who observed a temperature threshold of  $16^\circ\text{C}$  for spawning in the Bay of Saint-Brieuc.

## 2.5. Environmental forcing

### 2.5.1. Field data

The environmental variables used as forcing variables in the model are presented in Fig. 3. Three food proxies considered as the main food sources for scallops (e.g. Lorrain et al., 2002; Marchais, 2014) were monitored for our model in the Lanvéoc area (Fig. 1.): particulate organic matter (POM, expressed initially in  $\text{mg L}^{-1}$  but transformed a

posteriori to particles  $\text{L}^{-1}$ ), phytoplankton concentration (in cell  $\text{L}^{-1}$ ) and microphytobenthos concentration (also converted to cell  $\text{L}^{-1}$ ). Microphytobenthos concentration was provided by the IUEM (*Institut Universitaire Européen de la Mer*) observatory (Leynaert, 2016). As microphytobenthos concentration was not available for the studied period (1998–2003), we used a mean annual microphytobenthos profile that we applied for each year from 1998 to 2003 (Fig. 3b). POM data (in  $\text{mg L}^{-1}$ ) were transformed into the number of particles per litre considering each particle to have an average diameter of  $30\mu\text{m}$  (weight =  $1.4 \times 10^{-5}\text{ g}$ ; density = 1) (Lavaud et al., 2014).

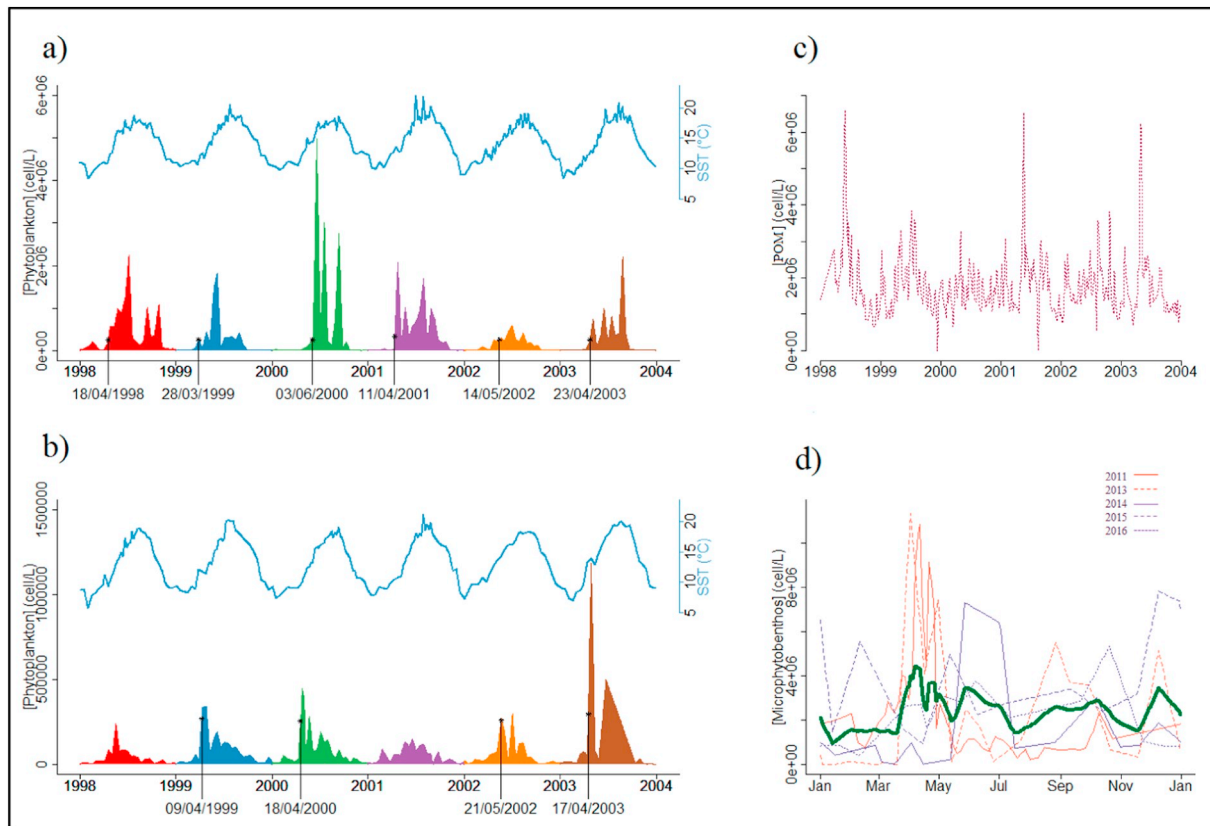
Weekly bottom temperatures were measured at Sainte-Anne from 1998 until 2003 by the SOMLIT monitoring network (Fig. 3a). Phytoplankton data and surface seawater temperature in the Bay of Saint-Brieuc from 1998 to 2003 were available from the REPHY monitoring network (Fig. 3c). All these environmental measurements were linearly interpolated to fit the daily time step of the simulations.

### 2.5.2. Climatic scenarios and forecasting approach

In order to study the potential effect of climate change on the reproductive cycle of *P. maximus*, we used monthly time series of predicted atmospheric temperature ( $T_{\text{Atm}}$ ) from the RCP2.6 (i.e., an increase of  $0.3$  to  $1.7^\circ\text{C}$ ) and RCP8.5 scenarios (a drastic increase of  $2.6$  to  $4.8^\circ\text{C}$ ) from 2040 to 2100 (IPCC Representative Concentration Pathways, Appendix A). For each scenario, we converted  $T_{\text{Atm}}$  into SST by using linear regressions according to the following equation:

$$\text{SST} = a \times T_{\text{Atm}} + b$$

where  $a$  is a coefficient that estimates the determination coefficient,

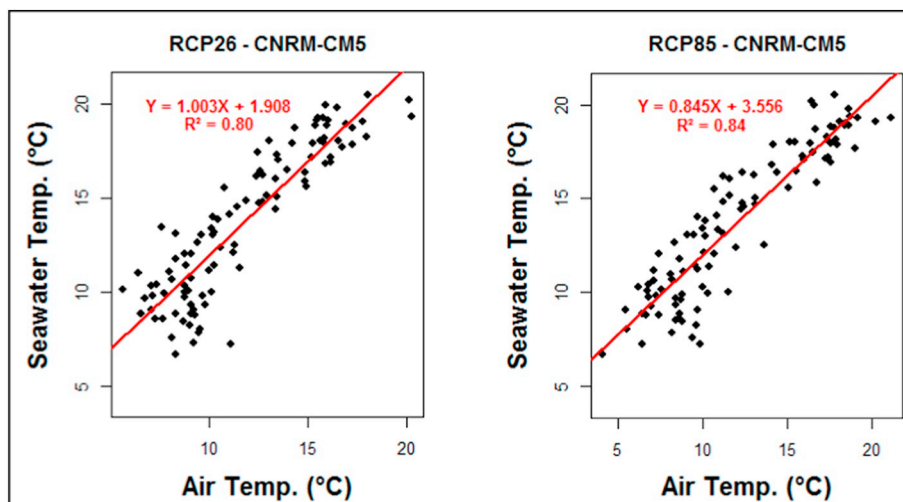


**Fig. 3.** Environmental forcing from 1998 to 2003 in (a, c, d) the Bay of Brest (Lanvéoc station) and (b) Bay of Saint-Brieuc (Saint-Pol station) used for the DEB model simulations. (a, b) phytoplankton concentration (one colour per year; cell/L) and seawater surface temperature (SST, light blue line; °C). The dates on x-axis indicate the time, each year, when the phytoplankton concentration threshold for spawning was reached ( $2.5010^5$  cell/L). (c) Particulate Organic Matter concentration (POM, magenta line; cell/L) measured in the Bay of Brest. (d) Microphytobenthos concentration measured in 2011 and from 2013 to 2016 in the Bay of Brest. Mean microphytobenthos concentration is shown by a green line. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

$b$  is a coefficient that represents the intercept (Fig. 4). To the best of our knowledge, no phytoplankton models or projections are currently available for our sites. Therefore, we used the previously recorded time series of phytoplankton, POM and microphytobenthos in the Bay of Brest from 1988 to 2003 (see previous paragraph) as potential conditions in future scenarios (letters A to F were used to refer to the phytoplankton concentrations observed

from 1998 to 2003, respectively).

We analysed patterns in reproductive activity in the simulations performed under the RCP2.6 and RCP8.5 scenarios. In each case,  $6 \times 20 = 120$  individual trajectories were simulated, with initial conditions corresponding to initial L and DFM of a representative set of individuals sampled during the 6-year monitoring program (see Scallop data in the results section).



**Fig. 4.** Relations between monthly air temperature from the RCP scenarios and monthly seawater temperature in the Bay of Brest (from 2006 to 2014): on the left, monthly air temperatures on monthly seawater temperature under the RCP2.6 scenario with the CNRM-CM5 model; on the right, monthly air temperatures on monthly seawater temperature under the RCP8.5 scenario with the CNRM-CM5 model.

## 2.6. Statistical analysis

To evaluate the best fit, mean simulations for each model (Mod-1, Mod-2A and Mod-2B) and mean observed data were compared using a Taylor diagram. This diagram provides a statistical summary of the agreement between a reference (observed data) and modelling results (Taylor, 2001). Three statistical measures are presented in the Taylor diagram: the centred root mean square (RMS) difference, normalized standard deviation, and Pearson's correlation coefficient. All statistical analyses were conducted in R version 3.3.3 (R Core Team, 2017).

## 3. Results

### 3.1. Contrasted environmental forcing conditions in the two study sites

Between 1998 and 2003, sea surface temperatures (SST) in the Bay of Brest reached a minimum of 8.3 °C in February 1998 and a maximum of 22.1 °C in July 2001 (Fig. 3a). The overall annual mean was  $14.0 \pm 0.3$  °C. The warmest year was 2001, with a yearly mean temperature of 14.5 °C. This year also had the warmest summer, with a mean temperature of 18.4 °C. The coldest year was 1998, with a yearly mean temperature of 13.7 °C. The coldest summer was 2000, with a mean temperature of 16.9 °C. Phytoplankton concentration from 1998 to 2003 averaged 328,000 cell/L per year, with an intra-annual SD of 135,000 cell/L. Phytoplankton concentration showed a seasonal pattern, with maximum values in spring and summer and minimum values in winter (Fig. 3a). The magnitude and timing of spring and summer blooms showed high inter-annual variability. For example, the spring bloom reached 5,000,000 cell/L in 2000, but the maximum phytoplankton concentration recorded in 2002 was 600,000 cell/L. The bloom onset date also differed among years. The first bloom observed in 2000 (30,000 cell/L) occurred on 20 January, while it was observed on 25 February in 1998 (206,000 cell/L).

The POM concentration showed similar patterns over the study period (Fig. 3a). However, larger peaks were observed in 1998, 2001 and 2003, at about 6,460,000 particles L<sup>-1</sup>, compared with lower values of 3,640,000 particles L<sup>-1</sup> in 1999, 2000 and 2002.

Microphytobenthos concentration showed two seasonal trends (Fig. 3b). The first pattern was observed in 2011 and 2013 with a large peak in spring and two smaller peaks in autumn and winter. The second pattern, observed in 2014, 2015, and 2016, showed a peak in early summer and two smaller ones in autumn and winter. The smallest number of microphytobenthic species ( $n = 22$ ) were identified in 2011 and a maximum of 67 species were identified in 2016. For the rest of the study, we used a mean profile of microphytobenthos computed by taking the average of all these observations (Fig. 3d).

In the Bay of Saint-Brieuc, SST fluctuated between a minimum of 5.8 °C in February 1998 and a maximum of 21.2 °C in July 2001 (Fig. 3b), thus showing a greater range of variation than the Bay of Brest. The average SST was  $12.6 \pm 0.5$  °C. The warmest year was 2003 with a mean temperature of 13.6 °C. This year also had the warmest summer with a mean temperature of 18.4 °C. As in the Bay of Brest, the coldest year was 1998, with a mean temperature of 12.6 °C, and 2000 was the coldest summer, with a mean temperature of 16.5 °C.

Phytoplankton concentrations were maximal in spring and summer and minimal in winter (Fig. 3b). The annual phytoplankton concentration from 1998 to 2003 averaged  $40,764 \pm 4990$  cell L<sup>-1</sup>. Bloom intensities were lower than in the Bay of Brest, but the magnitude and timing appeared quite different from year to year. For example, the 2003 spring bloom peaked at 1,187,000 cell L<sup>-1</sup>, while only reaching 150,000 cell L<sup>-1</sup> in 2001. The earliest first bloom was observed in 1999, on April 1 (114,000 cell L<sup>-1</sup>), while the latest was observed in 2001, on May 14 (103,000 cell L<sup>-1</sup>).

### 3.2. Comparing the DEB models

Simulations of dry flesh mass using model Mod-2A fitted the observations better than simulations from Mod-1 (Fig. 5). The addition of the fourth state variable  $E_{Go}$  seemed to improve prediction of spawning events. Spawning events can be identified on each curve by a sharp decline in DFM. The spawning period was more pronounced using Mod-2A than with Mod-1. For example, under Mod-1, the first spawning occurred on May 11 in 1999 whereas it appeared March 28 under Mod-2A. However, neither model successfully reproduced the observed increase in DFM from March to May. On average, the difference between observed and simulated DFM values from January to May was  $\pm 0.39$  g under Mod-1 and  $\pm 0.95$  g under Mod-2A. DFM modelled using Mod-2B was more accurate and the increase of DFM in spring fitted the observed data well ( $\pm 0.09$  g of difference). Similarly to Mod-2A, the spawning period was longer and more realistic than when using Mod-1. The addition of microphytobenthos to phytoplankton for *P. maximus* food intake allowed a better simulation of growth and reproductive activity, especially in the spring. For all years, model Mod-2B gave the best fit between observations and simulations of growth, with a mean correlation coefficient up to 0.9 and a normalized standard deviation close to 1 (Fig. 6). Therefore, for the remaining part of this study, we exclusively used model Mod-2B.

### 3.3. Simulation of reproductive activity

Simulations of individual DFM using Mod-2B in the two study areas highlighted three major trends in scallop reproductive activity. Firstly, the number of spawning events per reproductive season varies between years (Fig. 7). At least four spawning events were simulated for almost all individuals in the Bay of Brest except in 2000, when there were only three major spawning events. This could be because the phytoplankton threshold for spawning was only reached in June this year and the summer seawater temperatures were colder (Fig. 3).

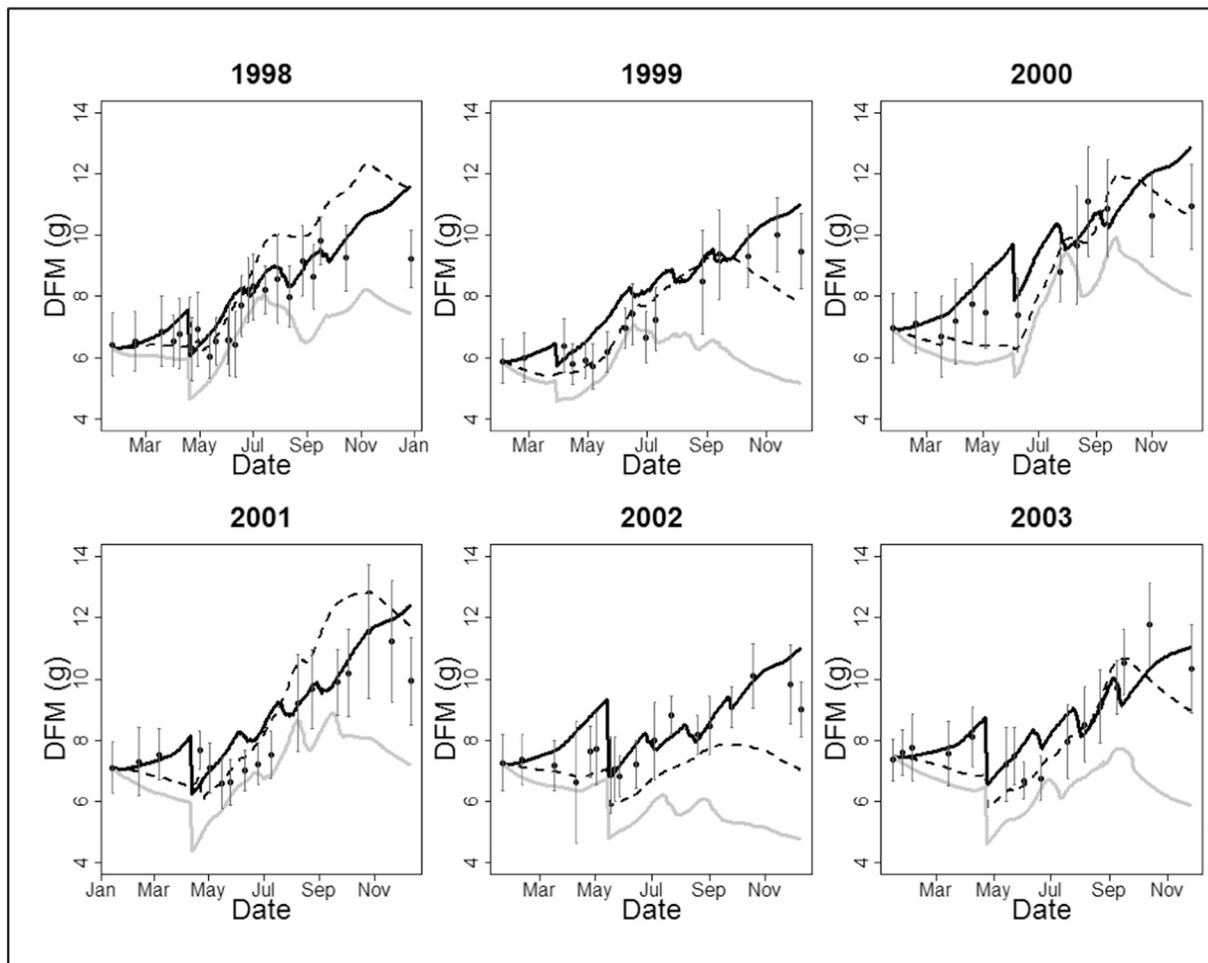
Secondly, the spawning period lasted from early spring to early autumn, corresponding to a wide period of 4–6 months depending on the year. This temporal window was shorter in 2000 and 2002 (around 100 days) compared with 1998 and 1999 (above 150 days; Fig. 7). The interval between spawning events (i.e. time for gametogenesis) ranged from 25 to 50 days. In the literature, spawning in *P. maximus* in the Bay of Brest was reported to span over a 6-month period with intervals of 20–50 days (Paulet et al., 1995). Our simulations are in full agreement with these field observations, showing the ability of the model to accurately simulate energy allocation to reproduction and spawning events in *P. maximus*.

The third observed pattern was asynchronicity between individuals, observed every year following the first synchronous spawning (except in 1999 when asynchronicity was also observed for the first spawning event). For instance, in 2003 spawning occurred in all individuals within 8 days while in 1999 it took 29 and 26 days for all the individuals to spawn during the first and fourth spawning events, respectively.

In comparison with results obtained for the Bay of Brest, only two clear spawning events were observed in the Bay of Saint-Brieuc (Fig. 8). Furthermore, in 2000, only 22% of individuals spawned twice. The spawning period was much shorter (~50 days) and only occurred in summer. The first spawning event was mainly synchronous between individuals, except in 1999. In contrast, the second spawning event was mainly asynchronous but the temporal window did not exceed 15 days.

### 3.4. Simulating reproductive activity of *P. maximus* in a warming ocean

In order to evaluate the impacts of climate change on the reproduction of *P. maximus* off the coast of Brittany, we simulated DFM in



**Fig. 5.** Mean observed ( $\pm$  SD,  $N = 20$ , black dots) and simulated ( $N = 20$ ) dry flesh mass (DFM) of 3-year-old scallops in the Bay of Brest between 1998 and 2003, using Mod-1 (dotted black line), Mod-2A (full grey line) and Mod-2B (full black line). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

two extreme cases: the RCP2.6 scenario (an increase of 0.3 to 1.7 °C) and the RCP8.5 scenario (a drastic increase of 2.6 to 4.8 °C) and six phytoplankton conditions (scenarios A to F).

In the Bay of Brest, under the RCP2.6 scenario, spawning dates were similar to current observations and no change in the individual spawning strategies were observed (Fig. 9a–b). Under the RCP8.5 scenario, we observed a decrease of the interval between spawning events after the first one except with phytoplankton regime D during the third spawning event (Fig. 9a–c). The second and third spawnings occurred 5 and 10 days earlier, respectively, under phytoplankton regimes C and E, and the fourth spawning occurred 20 days earlier under phytoplankton regimes B, D and F. Finally, more spawning occurred in autumn with a higher number of individuals spawning four to five times.

Spawning events were less frequent in the Bay of Saint-Brieuc than in the Bay of Brest in both temperature scenarios. Under the RCP2.6 scenario, two spawning events were observed as in the simulation of the current situation (Fig. 9d–e). In contrast, under RCP8.5 scenario, a third spawning event appeared every year (Fig. 9f). Moreover, spawning occurred 30 to 48 days earlier depending on the phytoplankton scenario. The spawning period was slightly shorter under the RCP2.6 scenario (around 40 days), while it was longer under the RCP8.5 scenario, reaching 80 days. Furthermore, asynchronicity between individuals was only observed under the RCP8.5 scenario, especially during the second spawning event.

Last, we found two significant relationships, between the date of the first spawning event and spring phytoplankton concentrations on one

hand and summer SST (mean value) on the other (Fig. 10). The date of the first spawning event was significantly inversely correlated with summer SST in the Bay of Saint-Brieuc (Fig. 10, linear regression:  $r^2 = 0.70$ ,  $p < 0.05$ , slope =  $-11 \text{ d } ^\circ\text{C}^{-1}$ ). In the Bay of Brest, the date of the first spawning event was significantly inversely correlated with the spring phytoplankton bloom concentration (Fig. 10, linear regression:  $r^2 = 0.74$ ,  $p < 0.05$ , slope =  $-1.22 \cdot 10^{-4} \text{ d cell L}^{-1}$ ).

#### 4. Discussion

The main objective of this study was to quantify the influence of environmental variables (temperature and food sources) on the reproductive processes of the great scallop, *P. maximus*, and explore the potential impacts of climate change on its dynamics. We improved an existing scallop DEB model developed by Lavaud et al. (2014) by detailing the reproductive processes and by adding microphytobenthos as a new food source.

In order to improve the DEB model for *P. maximus*, a fourth state variable was added to describe the fixation of energy in gametes, as done by Bernard et al. (2011) for the Pacific oyster, *Crassostrea gigas*. Furthermore, the maximum possible shell length was assumed to be 20 cm rather than the previous assumption of 12 cm, since studies have shown that adult scallops can reach 16 cm in the most favourable conditions (Mason, 1957; Chauvaud et al., 2012) and thus the ultimate length would presumably be above 16 cm. This led to the recalculation of three model parameters:  $\{p_{Am}\}$ ,  $[p_M]$  and  $\kappa$ . The new values obtained



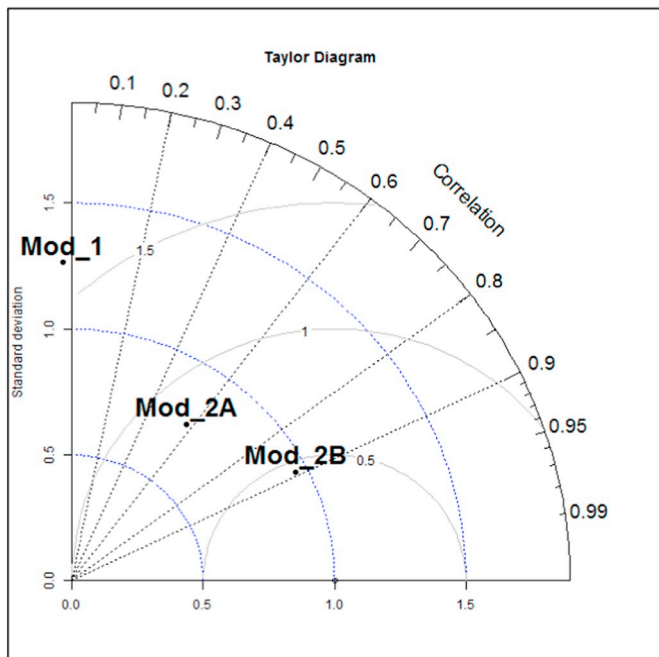


Fig. 6. Taylor diagram presenting the normalized standard deviation, Pearson's correlation coefficient and centred root mean squared difference (grey line) between simulated and observed dry flesh mass. The average of the 6 years simulated with each model is shown by a black dot.

are different from the previous version in Lavaud et al. (2014), particularly  $\kappa$ . The previous value, fixed at 0.86, was high compared with other bivalve species. For instance, the  $\kappa$  value for *Crassostrea gigas* is around 0.45. Considering that, in some environments, *P. maximus* could spawn more than three times within the same reproductive season and that its gonad represents > 30% of the whole flesh weight at maturity, it seems logical that this species would have a high energy allocation ratio (and thus a low value for  $\kappa$ ) similarly to *C. gigas*. Considering this, the new value calculated here is probably more consistent with the reproductive capacity of *P. maximus*. These changes do not fundamentally alter the dynamics of the model, but allow more spawning events and higher fecundity than other versions of the model. Of course, further testing in other locations with contrasted forcing conditions as well as with younger age-classes would also be needed to fully validate this updated version of the scallop DEB model.

Another improvement made in the current model concerns trophic resources. Microphytobenthos was added as a new source of food for scallops. Previously, Lavaud et al. (2014) demonstrated that POM constitutes an additional food source allowing scallops to compensate phytoplankton limitation. In addition, our study suggests that microphytobenthos would probably also be a valuable source of food that could sustain energy acquisition, especially in spring. For the moment, the taxonomic composition of each food source is not detailed in the model, but several studies have shown relationships between specific phytoplankton species and life history traits of the great scallop in Brittany. For example, Chauvaud et al. (1998, 2001) showed, in the Bay of Brest, that growth and food intake of *P. maximus* were dependent on phytoplankton taxonomic composition and concentration. The related growth cessation depended on massive sedimentation of diatom blooms or toxic dinoflagellate blooms. For example, *P. maximus* food intake and growth were highest when *Cerataulina pelagica* blooms occurred and lowest during *Gymnodinium nagasakiense* blooms. In addition, Lorrain et al. (2000) demonstrated that large bottom concentrations of chlorophyll-*a*, following diatom blooms, could have a negative effect on

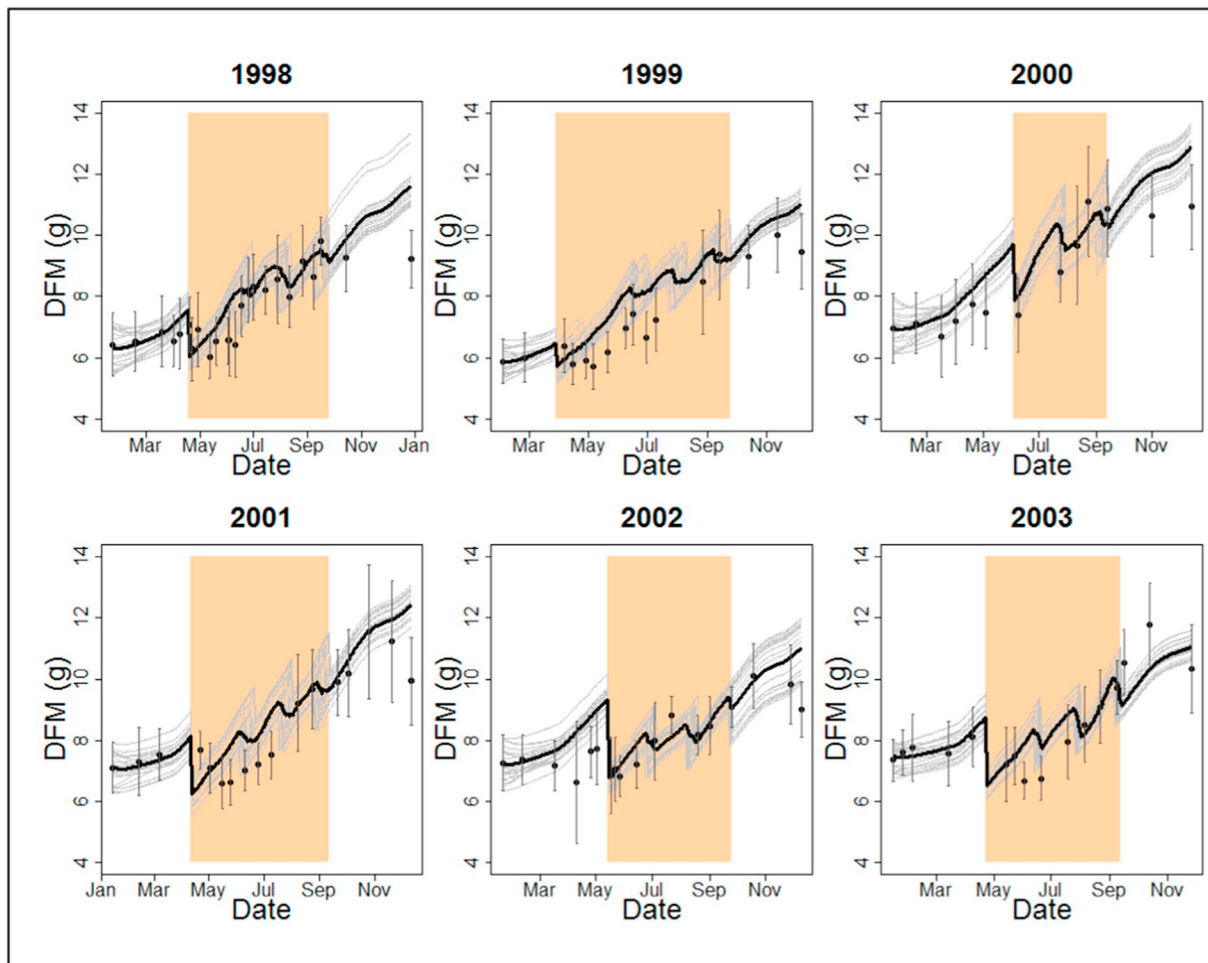
the ingestion or respiration of *P. maximus* juveniles, either by gill clogging or by oxygen depletion at the water-sediment interface associated with the degradation of organic matter. The current version of the DEB model does not take into account these specific effects which are linked to the type of food that is available. However the current model provides the basis for taking them into consideration in future studies.

One major difficulty with a modelling approach is to obtain a sufficient dataset to calibrate, test and validate a numerical model. When using a bioenergetics model, this implies monitoring growth and reproduction of marine organisms and their surrounding environmental data, at the same place and ideally over a long period (many years) to evaluate temporal phenotypic variability. In our study, the Bay of Brest sampling sites (St Anne, Roscanvel, Lanvéoc) are not closed off from each other but are instead located in a very well mixed area within the Bay of Brest (Salomon and Breton, 1991) where scallops are the most abundant. So we can suppose that environmental data are sufficiently representative of conditions encountered by scallops. In the Bay of Saint-Brieuc, there is no growth monitoring of scallops and there are too many gaps in the environmental data to apply the model in a satisfactory manner. Our approach is therefore limited but it offers a first application of the model to this new environment and constitutes a stepping-stone for further development of the modelling approach for this bay.

Other limitations of our model that should be mentioned are its systematic overestimation of growth during the autumnal period and an insufficient integration of inter-individual variability. The systematic tendency to overestimate growth could be due to a change in the physiology of scallops at the end of the reproductive season and period leading into winter. Specific ecophysiological experiments should be developed to address this question and improve the model. For the moment, we have applied an individual-based modelling strategy by introducing variability between individuals through the initial condition values. To account for more variability in physiological traits, similar studies, e.g. Thomas et al. (2016) and Bacher and Gangnery (2006) used specific model parameterization of the ingestion function for each individual. For instance,  $X_k$  values were allocated to each individual following a Gaussian distribution. It would now be interesting to adapt a similar approach to the scallop DEB model.

Quantitative modelling of reproductive processes (preliminary storage phase, gametogenesis, spawning and/or resorption) is not easy as these processes are typically species-specific. There are no general rules on how to handle reproductive activity in DEB theory, especially regarding reproduction buffer dynamics. Bernard et al. (2011) introduced a fourth state variable in order to improve modelling of reproduction dynamics in the Pacific oyster, *Crassostrea gigas*. Numerous marine organisms from temperate waters spawn once or twice at a relatively fixed time each year (Gosling, 2004). For *P. maximus*, however, reproductive activity is more complex, with asynchronous spawning during a highly variable reproductive window. For this preliminary approach, however, we made the assumption that the mechanisms governing reproductive activity would be quite similar among bivalves and thus between oysters and scallops.

The reproductive cycle of *P. maximus* has been studied extensively in many places (e.g. Magnesen and Christophersen, 2008). Concerning our studied area, contrasting patterns were shown between the Bay of Brest and Bay of Saint-Brieuc (Paulet et al., 1995). Scallops in the Bay of Brest usually spawn between April and October, with a massive first spawning in April followed by a second maturation phase until July characterized by one or several smaller spawning events. A third summer maturation phase leads to a last major spawning event during August (Paulet et al., 1995). A late spawning event may be observed in autumn, but only in a few individuals (Saout, 2000). In the Bay of Brest, this seasonal cycle varies



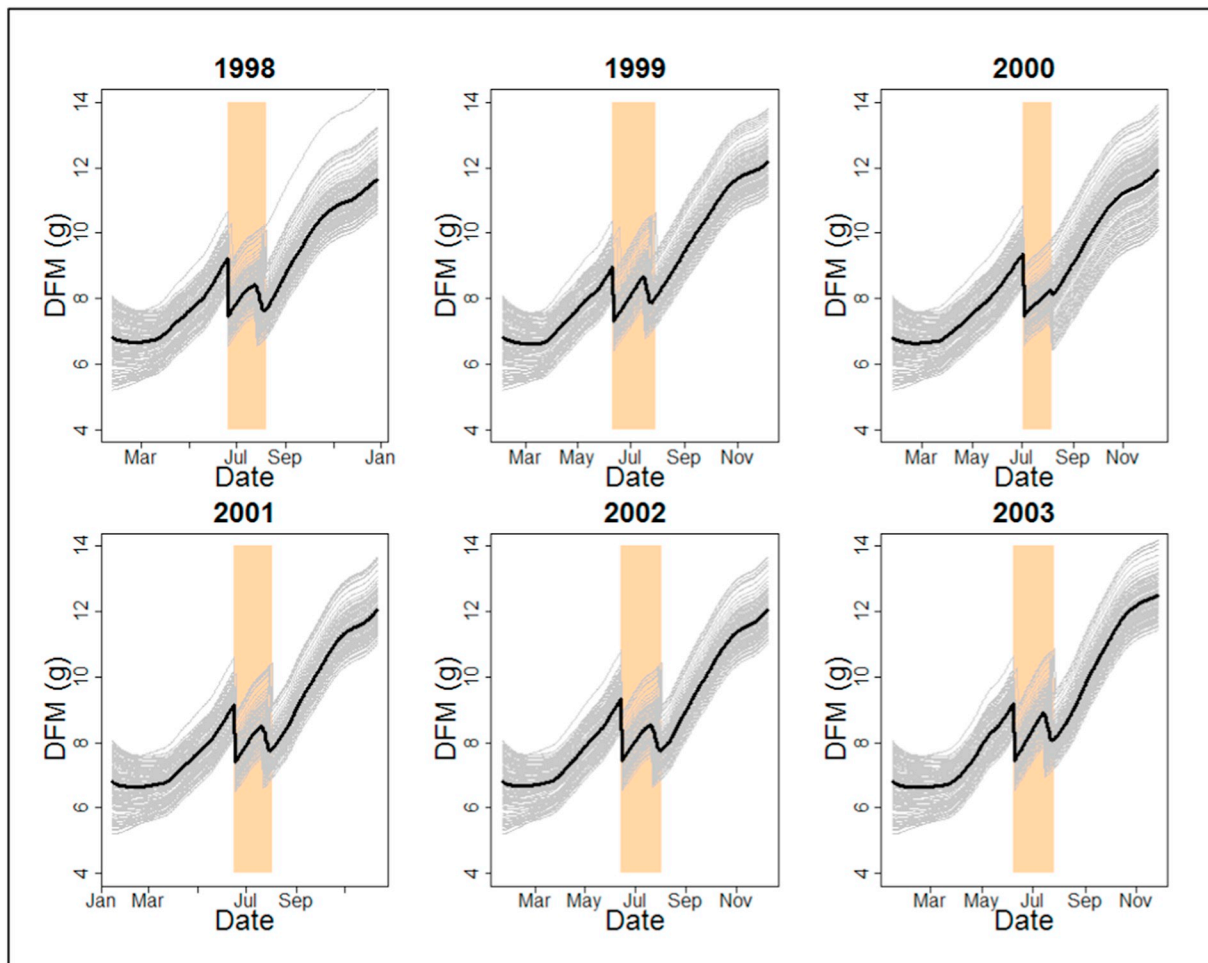
**Fig. 7.** Mean observed ( $\pm$  SD,  $N = 20$ , black dots) and simulated ( $N = 20$ , thick black line) dry flesh mass (DFM) of 3-year-old scallops in the Bay of Brest between 1998 and 2003 using Mod-2B. Individual growth trajectories of the 20 scallops simulated are indicated by grey lines. The orange area shows the spawning period. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

strongly between individuals, resulting in a lack of synchronism (Paulet et al., 1995). The results from our simulations, obtained over six years in the Bay of Brest, correspond fairly well to this description. More precisely, the model was able to partly simulate the observed asynchronicity between individuals, and the mechanisms implemented to trigger spawning appeared to be sufficiently relevant to simulate the onset of the first spawning, the temporal spawning window, and the number of spawning events (Fig. 6).

In the Bay of Saint-Brieuc, scallops only spawn from June until mid-August (Paulet et al., 1988; Paulet, 1990). Moreover, the seasonal cycle is known to be similar between individuals, showing a higher synchronism in this area than in the Bay of Brest (Paulet et al., 1988). Of course, the application of our model in the Bay of Saint-Brieuc is only a first attempt and suffers from a lack of forcing data. Nevertheless, it seems that the current version of the model was able to reproduce the different patterns of the reproductive cycle of *P. maximus* observed in this area. This tends to confirm that the environmental factors used here are the main key-drivers of reproduction processes of *P. maximus*.

Despite its limitations, our modelling study suggests that differences in the timing of spawning events might be explained mainly by environmental differences in food and temperature. Among invertebrates, there is much evidence for the influence of exogenous factors (e.g. temperature, food and photoperiod) on the progress of gametogenesis and for regulation by endogenous rhythms on which environmental signals may act as synchronizers (e.g. Mat et al., 2014). Many

environmental variables trigger such regulation but, most often, temperature and food availability (Franco et al., 2015; Ubertaini et al., 2017) are considered to be the key factors. This is the case for bivalves, particularly pectinids (Lavaud et al., 2014). Scallops are sublittoral, epifaunal and active suspension-feeding bivalves that rely on suspended detrital material, phytoplankton and microphytobenthos as their main food sources (Robert et al., 1994; Chauvaud et al., 2001). Saout et al. (1999) showed that, in *P. maximus*, a simultaneous increase of temperature and photoperiod enhanced gonad growth when food is not limiting. However, it was still not clear which of these two factors is the more influential. Our results obtained in the Bay of Brest show that, within a temporal photoperiod window, spawning is mainly triggered mainly by phytoplankton blooms once the GSI threshold is reached. In this eutrophic area, temperature might play a secondary role in terms of triggering spawning. For instance, in 2000 and 2002, the first bloom of the year was late compared with the other years (June 3 and May 14, respectively; Fig. 3). Accordingly, for both years, the model simulated a later occurrence of first spawning (June 4 and May 15, respectively; Fig. 7) that fitted well with field observations. By comparison, blooms of phytoplankton in the Bay of Saint-Brieuc were much lower than those observed in the Bay of Brest and presumably not a source of stress. Values were always below the threshold for triggering spawning. In this new environment, phytoplankton blooms were presumably not the trigger for spawning. Based on previous studies, we basically used a temperature threshold in this environment (Fifas, 2004) and the



**Fig. 8.** Individual (grey lines) and mean simulated ( $N = 120$ , thick black line) dry flesh mass (DFM) of 3-year-old scallops in the Bay of Saint-Brieuc in 1998–2003 using Mod-2B. The orange area shows the spawning period. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

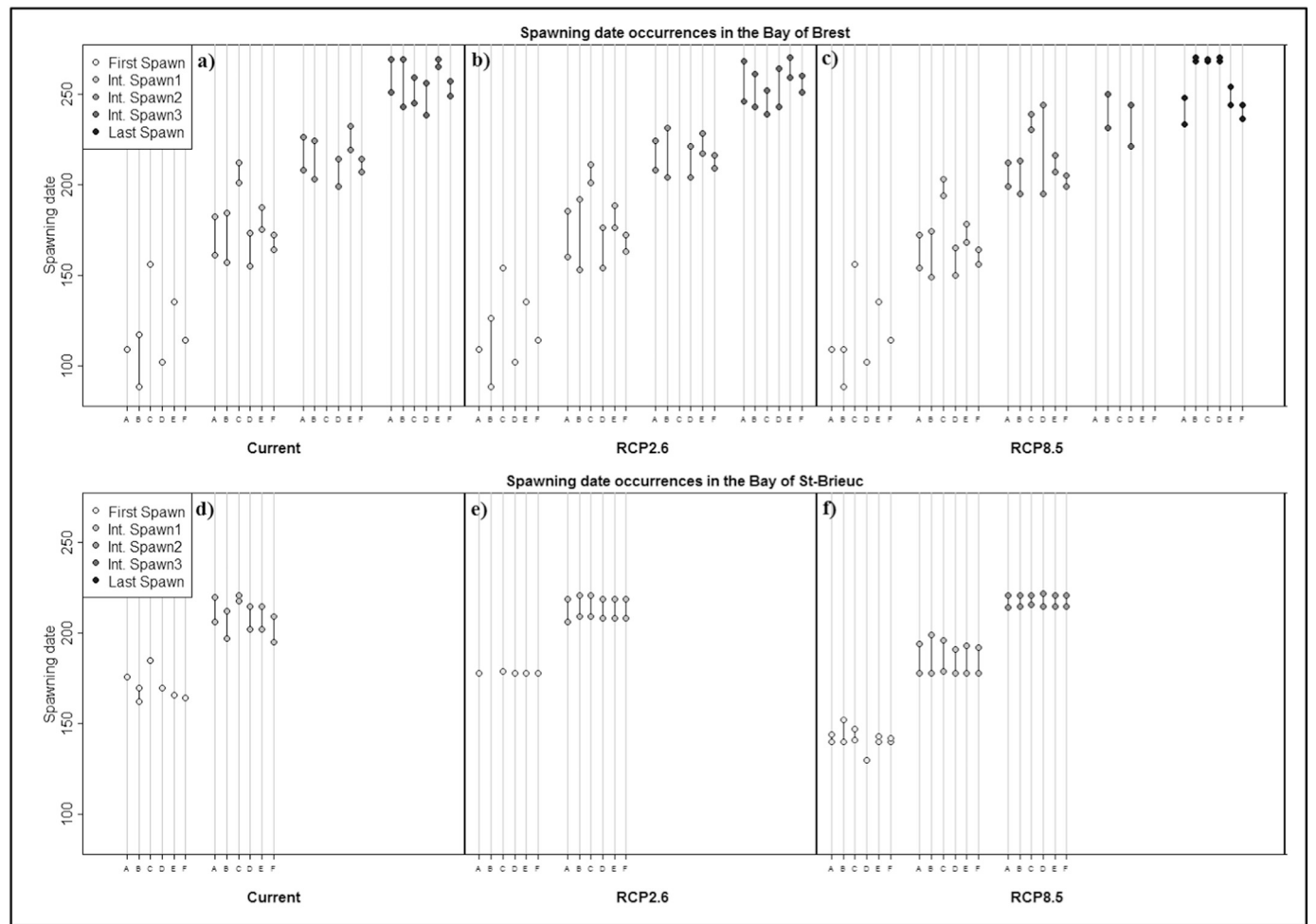
simulations obtained were in accordance with spawning processes observed in this bay.

In the last part of this study, we explored the potential consequences of climate change for the reproductive activity of *P. maximus* in northern Brittany. Predicting the temperature of the future atmosphere and oceans has been a focus of research for a few decades now. The evolution of food supply to organisms, which in the ocean starts with phytoplankton, is comparatively less well understood or predictable. Although the reliability of ocean primary production models is continually improving, predicting the future is challenging (Gradinger, 2009; Lavoie et al., 2017) for coastal environments. In this context, we believe that our approach, consisting of transposing current food availability time series to future scenarios, is valuable because it allows a greater degree of complexity in predictions as it provides realistic estimates of inter-annual variability. This approach could be complemented by simulations under enhanced phytoplankton productivity, as predicted by recent modelling (Jensen et al., 2017). Only the most extreme scenario (RCP8.5) revealed contrasting predictions with the current one. While distinct reproductive cycles are currently observed between the Bay of Brest and the Bay of Saint-Brieuc, it seems that future environmental conditions would generally extend the spawning period, with an additional spawning event predicted in both locations. However, contrasted impacts also emerged when comparing

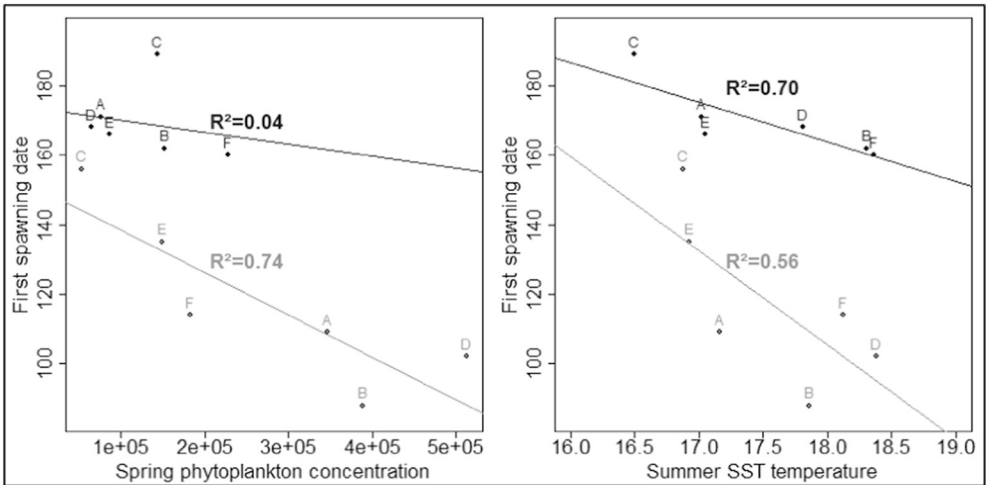
simulations obtained for the two bays. An increase in seawater temperature advanced the onset of spawning by 20–44 days in the Bay of Saint-Brieuc, irrespective of the phytoplankton scenario, while the spawning timeline in the Bay of Brest was unchanged. Warmer temperatures might also lead to better recruitment. Shephard et al. (2010) found that the mean annual recruitment of young scallops in the Isle of Man was positively related to spring water temperature. Adult gonads were also larger, indicating higher egg production during warmer years. Our simulations led to similar conclusions, showing that an increase in seawater temperature combined with adequate food supply could well enhance scallop recruitment by: (1) increasing the spawning window in late summer and (2) advancing the onset of spawning in spring in the Bay of Saint-Brieuc.

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**Fig. 9.** Simulated spawning date (Int. = Intermediate) in the Bay of Brest (a to c) and Bay of Saint-Brieuc (d to f) under three temperature scenarios (current, RCP2.6 and RCP8.5) and six phytoplankton regimes (A to F, corresponding to conditions in years 1998 to 2003). The lines between two points represent the asynchronicity between individuals with the first and last spawning date within a population.



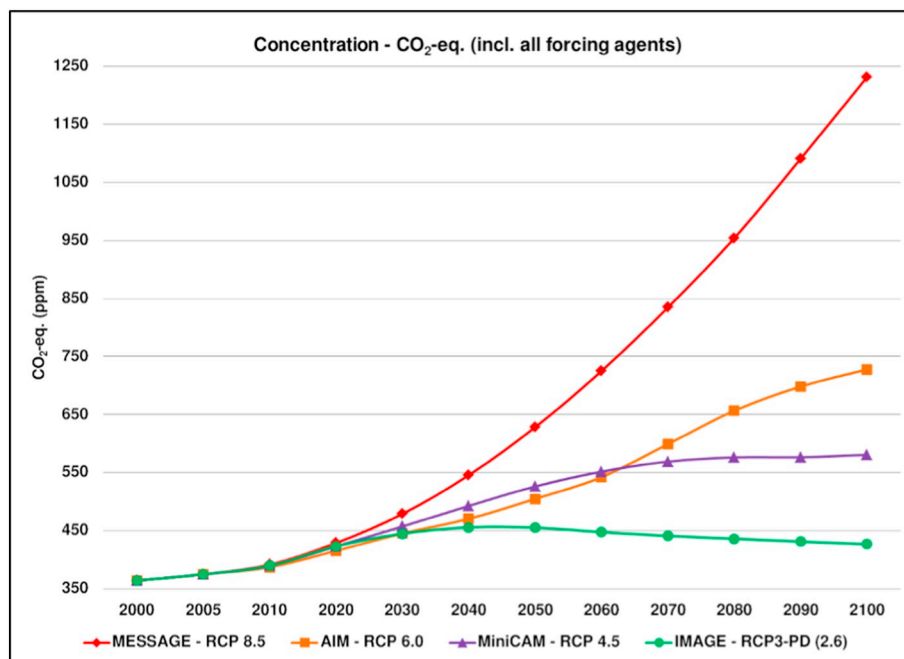
**Fig. 10.** Relationships between environmental conditions (left: spring phytoplankton concentration; right: summer SST temperature) and the first spawning date in the Bay of Brest (grey) and in Bay of Saint-Brieuc (black) for the six phytoplankton regimes (A to F). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

**Appendix A. Details on climatic projections models used and additional figures (Figs. A1 and A2) and tables (Table A1 and A2)**

In this study we used temperature scenarios resulting from Representative Concentration Pathways (RCP) models, the latest generation of scenarios that provide inputs to climate models. The purpose of using scenarios is not to predict the future, but to explore both the scientific and real-world implications of different plausible futures (van Vuuren et al., 2011). The IPCC authors chose four carbon dioxide ( $\text{CO}_2$ ) emission trajectories to focus on and labeled them based on how much heating they would result in at the end of the century: 2.6, 4.5, 6 and 8.5 watts per square meter



( $\text{W m}^{-2}$ ). Fig. A1 shows the annual  $\text{CO}_2$  emissions (in billions of tons of carbon) until 2100 for each of the RCPs.



**Fig. A1.** Emissions of annual  $\text{CO}_2$  across the RCPs. The RCP2.6 scenario (IMAGE-RCP3-PD(2.6)) and the RCP8.5 scenario (MESSAGE-RCP8.5) represent extreme situations: RCP2.6 is the most optimistic and RCP8.5 is the most drastic warming. Source: [IPCC Fifth Assessment Report \(2014\)](#).

In the  $2.6 \text{ W m}^{-2}$  scenario (RCP2.6), greenhouse-gas emissions drop to zero by about 2070, and then continuing to fall, so that the world's emissions would become negative — actually withdrawing greenhouse gases from the air and locking them away — for decades. This pushes the bounds of what is plausible through mitigation, some experts say. At the high end, in the  $8.5 \text{ W m}^{-2}$  scenario (RCP8.5),  $\text{CO}_2$  levels would soar beyond 1300 ppm by the end of the century and continue to rise rapidly (Table A1).

Table A1

Description of  $\text{CO}_2$  emissions scenarios used by IPCC authors ([van Vuuren et al., 2011](#)).

Scenario	Description
RCP8.5	Rising radiative forcing pathway leading to $8.5 \text{ W m}^{-2}$ ( $\sim 1370 \text{ ppm CO}_2 \text{ eq}$ ) by 2100.
RCP6	Stabilization without overshoot pathway to $6 \text{ W m}^{-2}$ ( $\sim 850 \text{ ppm CO}_2 \text{ eq}$ ) at stabilization after 2100.
RCP4.5	Stabilization without overshoot pathway to $4.5 \text{ W m}^{-2}$ ( $\sim 650 \text{ ppm CO}_2 \text{ eq}$ ) at stabilization after 2100.
RCP2.6	Peak in radiative forcing at $\sim 3 \text{ W m}^{-2}$ ( $\sim 490 \text{ ppm CO}_2 \text{ eq}$ ) before 2100 and then decline (the selected pathway declines to $2.6 \text{ W m}^{-2}$ by 2100).

Atmospheric temperature data were obtained from the CERFACS modelling center. For each scenario (RCP2.6 and RCP8.5) 14 models were available (<http://cmip-pcmdi.llnl.gov/cmip5/availability.html>; Table A2). To know which model was the most comparable to our historical temperature data in the Bay of Brest and the Bay of Saint Brieuc, we used the diagram of Taylor (Fig. A2) in order to compare monthly air temperature since 1960 to nowadays in our bays with monthly air temperature from the 14 models during the same period. Among the 14 models, the CNRM-CM5 model was the best (Fig. A2).

Table A2

Description of the 14 models available for the study.

Modelling center	Model	N°	Institution
BCC	BCC-CSM1.1	1	Beijing Climate Center
CCCma	CanESM2	2	Canadian Centre for Climate Modelling and Analysis
NCAR	CCSM4	3	National Center for Atmospheric Research
NSF-DOE-NCAR	CESM1(CAM5)	4	National Science Foundation, Department of Energy, National Center for Atmospheric Research
CNRM-CERFACS	CNRM-CM5	5	Centre National de Recherches Météorologiques / Centre Européen de Recherche et Formation Avancées en Calcul Scientifique
LASG-IAP	FGOALS-g2	6	LASG, Institute of Atmospheric Physics
NOAA GFDL	GFDL-CM3	7	Geophysical Fluid Dynamics Laboratory

(continued on next page)

Table A2 (continued)

Modelling center	Model	N°	Institution
INPE	HadGEM2-ES	8	Instituto Nacional de Pesquisas Espaciais
IPSL	IPSL-CM5A-LR	9	Institut Pierre-Simon Laplace
	IPSL-CM5A-MR	10	
MPI-M	MPI-ESM-LR	11	Max Planck Institute for Meteorology
	MPI-ESM-MR	12	
MRI	MRI-CGCM3	13	Meteorological Research Institute
NCC	NorESM1-M	14	Norwegian Climate Centre

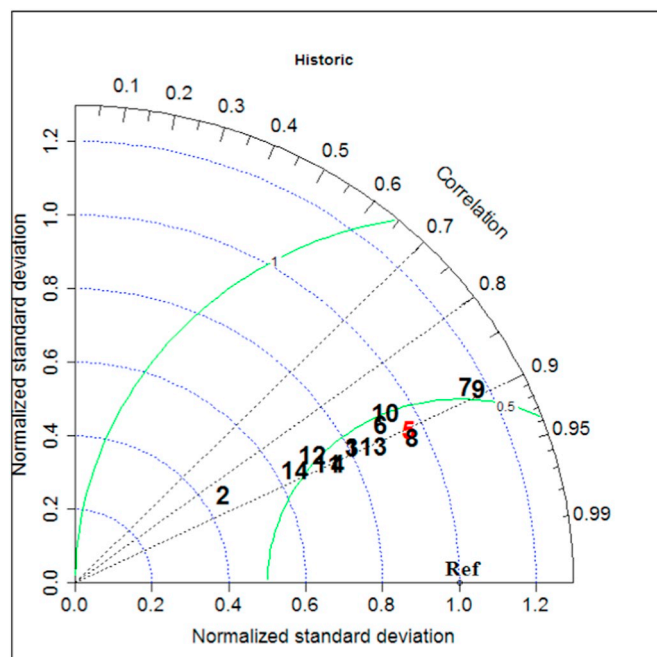


Fig. A2. Taylor diagram giving the Pearson's correlation coefficient, the centered root mean squared difference and the normalized standard deviation between historical air temperature observed in the Bay of Brest since 1960 (Ref) and air temperature predicted by 14 different models (see Table above; N°1 to N°14).

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